# BB-SPEEDset: A Validated Dataset of Broadband Near-Source Earthquake Ground Motions from 3D Physics-Based Numerical Simulations

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#### ABSTRACT

This article introduces a strong-motion dataset of near-source broadband earthquake ground motions from 3D physics-based numerical simulations-named BB-SPEEDsetobtained by the code SPEED (SPectral Elements in Elastodynamics with Discontinuous Galerkin)—developed at Politecnico di Milano, Italy. Taking advantage of the earthquake ground-motion scenarios produced so far by SPEED, in most cases validated against earthguake recordings, the main objective of this work is to construct and validate a dataset of simulated broadband waveforms to be used as a support for characterization and modeling of near-source earthquake ground motions. To pursue this objective, the following steps were necessary, namely: (1) the implementation of an effective workflow suitable to process in an homogeneous format various SPEED simulations; (2) the generation of broadband time histories using a technique based on artificial neural networks, trained on strong-motion records; (3) the creation of a flat file collecting, for each simulated scenario, the most relevant metadata (fault rupture scenario, site response proxies, source-to-site distances) as well as a comprehensive set of ground-motion intensity measures of the processed broadband waveforms (peak ground acceleration, velocity and displacement, spectral ordinates, duration, pulse period, etc.). Finally, a comprehensive set of consistency checks is made to verify the absence of any systematic bias in the trend of the BB-SPEEDset results with respect to the NEar-Source Strong-motion (NESS) version 2.0 near-source recorded ground-motion dataset. Indeed, the main features of near-source ground motion in BB-SPEEDset, ranging from the statistical distributions of peak and integral measures both at short and long periods, the ground-motion attenuation with distance, to the features of impulsive ground motions and directionality effects, are in substantial agreement with those from NESS.

#### **KEY POINTS**

- A dataset of broadband near-source ground motions from 3D physics-based numerical simulations is created.
- The features of BB-SPEEDset are consistent with the NEar-Source Strong-motion dataset (NESS).
- BB-SPEEDset may serve as a support for an improved characterization of near-source ground motions.

### INTRODUCTION

It is well known that the characterization of earthquake ground motion in the near-source region is made difficult by the paucity of records that, in spite of their evergrowing number, cannot reliably describe yet neither the median values nor their variability, in the variety of source and site conditions typically present in the vicinity of the seismogenic fault.

As in most fields of science, when the laboratory investigations are either limited or prevented owing to the size of the prototype and to the difficulties to reproduce the in-field conditions, analytical and numerical modeling may be an alternative to complement in an ideal laboratory the information that is difficult to capture from nature.

In this perspective, the so-called physics-based numerical simulations (PBS) of earthquake ground motion aim at complementing the recorded data by providing simulated results in the source and site configurations that may resemble as closely as possible the real ones. In some cases, the role of PBS has been extended to provide realistic seismic scenarios of earthquake ground motions suitable to improve the approaches for

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**Cite this article as** Paolucci, R., C. Smerzini, and M. Vanini (2021). BB-SPEEDset: A Validated Dataset of Broadband Near-Source Earthquake Ground Motions from 3D Physics-Based Numerical Simulations, *Bull. Seismol. Soc. Am.* **XX**, 1–19, doi: 10.1785/0120210089

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seismic hazard and risk analysis (see e.g., Graves *et al.*, 2011; Maeda *et al.*, 2016; Bradley *et al.*, 2017, 2020; Smerzini and Pitilakis, 2018; Stupazzini *et al.*, 2021) and to provide input for seismic structural analyses (Galasso *et al.*, 2013; Baker *et al.*, 2021; Fayaz *et al.*, 2021).

With this objective, several research groups worldwide (see e.g., amongst others, Graves and Pitarka, 2010, 2015; Irikura and Miyake, 2011; Komatitsch *et al.*, 2013; Mazzieri *et al.*, 2013; Isbiliroglu *et al.*, 2015; Lu *et al.*, 2018; Paolucci *et al.*, 2018; McCallen, Petersson, *et al.*, 2020; McCallen, Petrone, *et al.*, 2020) have continuously contributed in the recent years to the development of numerical tools that may become more and more suitable to produce, with a reasonable computational effort, realistic earthquake ground motions that may reliably complement the recorded ones in the near-source region and eventually be coupled with engineering models for non-linear structural response.

Combining the computational burden with the difficulty to accurately reproduce details of the fault geometry, of the seismic slip distribution, and of the complex geology (typically 3D) of the area of interest, that may extend by tens of kilometers, PBS are generally considered to be bounded within frequency limits hardly beyond about 2–3 Hz, although some successful examples of PBS extending up to 8–10 Hz in the presence of very detailed knowledge of local site conditions are also present, such as for the simulation of induced seismicity in the Groningen (the Netherlands) area (Paolucci *et al.*, 2021).

Some cross-verification activities of numerical tools for PBS were undertaken in the recent past (Bielak et al., 2010; Chaljub et al., 2010; Maufroy et al., 2015) that were seminal steps for the different research groups to solve the major issues arising when the numerical codes are applied to very complex configurations. However, relatively little effort was devoted up to now to comprehensive validations of PBS against strong-motion records, especially in the near-source region (Taborda and Bielak, 2013, 2014; Paolucci et al., 2015, 2021; Imperatori and Gallovič, 2017; Gatti et al., 2018; Pitarka et al., 2020). For this reason, a blind prediction experiment was set up in the framework of the 6th International Symposium on the Effects of Surface Geology on Seismic Motion (ESG6), jointly supported by the International Association of Seismology and Physics of the Earth's Interior (IASPEI) and the International Association of Earthquake Engineering (IAEE), with the objective to reproduce earthquake ground motions during the Kumamoto, Japan, seismic sequence of 2016, with a moment magnitude  $M_w$  7 mainshock.

With a long-lasting expertise gained (1) in the development of the open-source numerical code SPEED based on spectral elements (Mazzieri *et al.*, 2013), (2) in the advancement of techniques to enrich at high frequencies the PBS results (Paolucci *et al.*, 2018), (3) in the validation of PBS results against nearsource ground motions recorded from different earthquakes in Italy and worldwide, and in the application to several scenario case studies (see overview in Table 1), we have collected a large subset of our simulated results with a uniform processing procedure that will be illustrated in the sequel. In this way, we have constructed the BB-SPEEDset version 1.0 (v.1.0)-a dataset of broadband near-source ground motions aiming at providing a complementary tool for the characterization of earthquake ground motions, in terms of their dependency on magnitude, distance, and site conditions, such as the most common empirical ground-motion models (GMMs), with a complete and well-constrained description in terms of seismic source and site conditions. In addition, this dataset may also provide the basis to properly analyze the spatial variability of ground motion, with potential important implications to validate and improve the existing models for spatial correlation (e.g., Infantino, Smerzini, and Lin, 2021; Schiappapietra and Smerzini, 2021) and spatial coherency that currently suffer from the lack of records from sufficiently densely spaced arrays of seismic stations (e.g., Smerzini, 2018).

As introduced by D'Amico *et al.* (2017), the main advantage of a broadband ground-motion dataset based on PBS is that all input source data are clearly identified, as well as the site conditions of recording stations, and ground-shaking scenarios of the simulated earthquakes can easily be constructed. On the other side, it is difficult to prove that the available waveforms are not biased with respect to records, in terms of the different parameters of ground motion that are relevant for engineering applications, typically because of the limited detail of the input data of PBS in terms of seismic source and geological layering, and because of the computational limits of the numerical simulations.

To overcome such limitations, the novelty of this article is to provide a comprehensive comparison of the BB-SPEEDset statistical distributions with those obtained by the NEar-Source Strong-motion dataset NESS (v.2.0)-a dataset of worldwide recorded near-source ground motions addressed in this special issue by Sgobba et al. (2021), updated after Pacor et al. (2018). Although the origin of the two datasets is completely different, BB-SPEEDset being based on relatively few earthquake scenarios, each with a large sample of simulated accelerograms, whereas NESS is based on a large number of real earthquakes, each with relatively few records, such comparison is a crucial step to assess whether a bias exists between the trend of simulated results in the BB-SPEEDset with respect to those of NESS. The absence of systematic differences will strongly support the effectiveness of the procedure to produce broadband waveforms from the PBS as well as the potential use of the BB-SPEEDset to improve the available tools for the prediction of near-source earthquake ground motion and to provide input motions for earthquake engineering analyses.

This article is organized as follows. After an introduction of the workflow for postprocessing PBS results and for generating broadband waveforms by taking advantage of artificial neural networks (ANNs), the BB-SPEEDset is introduced with its current features, in terms of distribution of simulated waveforms according to magnitude, distance, and site conditions. Subsequently, a comprehensive comparison of the

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Case Study	Fault (SoF)	М	Model Size (km³)	V <sub>S,min</sub> (m/s)	f <sub>max</sub> (Hz)	Event for Validation	References	Included in BB-SPEEDset
Grenoble, France Gubbio plain, central Italv	Belledonne (SS) Colfiorito (NF)	6.0	41 × 50 × 8 85 × 62 × 10	300 250	რო	Benchmark 1997/09/26	Stupazzini e <i>t al.</i> (2009) Smerzini et <i>al.</i> (2011)	Not included yet Not included yet
Tagliamento plain,	Gemona Faults (TF)	6.1	57 × 53 × 12	300	2.5	1976/09/15	Smerzini (2010)	Not included yet
L'Aquila, central Italy	Paganica (NF)	6.2	58 × 58 × 20	300	2.0	2009/04/06	Evangelista <i>et al.</i> (2017)	2009 6.2 Aquila
Sulmona, central Italy	Mt. Morrone (NF)	6.0 6.5	49 × 42 × 13	500	2.5	Ideal scenarios	Villani <i>et al.</i> (2014)	6.0 SulmonaS03, 6.0 SulmonaS04 6.5 SulmonaS03, 6.5 SulmonaS05
Christchurch, New Zealand	Lyttelton (TF)	6.3	60 × 60 × 20	300	2.0	2011/02/22	Guidotti <i>et al.</i> (2011)	Not included yet
Po Plain, northern Italy	Mirandola (TF)	6.0	$74 \times 51 \times 20$	300	1.5	2012/05/29	Paolucci <i>et al.</i> (2015)	2012 6.0 Emilia
Marsica, central Italy	Fucino (NF)	6.7	56 × 46 × 20	100	2.0	1915/01/13	Paolucci <i>et al.</i> (2016)	1915 6.7 Marsica
Thessaloniki, northern Greece	Gerakarou (NF)	6.5	82 × 64 × 31	300	1.5	1978/06/20	Smerzini <i>et al.</i> (2017)	1978 6.5 Salonicco
	Anthemountas (NF)	7.0				Ideal scenario	Smerzini <i>et al.</i> (2018)	7.0 SaloniccoS01
Norcia, central Italy	Mt. Vettore–Mt. Bove (NF)	6.5	$50 \times 40 \times 21$	280	1.5	2016/10/30	Özcebe <i>et al.</i> (2019)	2016 6.5 Norcia
		5.8				Ideal scenarios		5.8 NorciaS01 5.5 NorciaS01
Wellington, New Zealand	Wellington-Hutt (SS)	6.0-7.0	80 × 50 × 45	300	2.0	Ideal scenarios	Paolucci <i>et al.</i> (2014)	Not included yet
Santiago, Chile	San Ramon (TF)	6.0, 6.5, 7.0	97 × 77 × 19	400	2.0	2010/04/01	Pilz <i>et al.</i> (2011)	Not included yet
Istanbul, Turkey	North Anatolian fault– Marmara Sea (SS)	5.7	165 × 100 × 30	250	1.5	2019/09/26	Infantino <i>et al.</i> (2021)	Not included yet
		7.0				Ideal scenarios	Stupazzini <i>et al.</i> (2021)	7.0 IstanbulS16, 7.0 IstanbulS20,
		7.2				ldeal scenarios		7.2 IstanbulS05, 7.2 IstanbulS09,
		7.4				Ideal scenarios		7.2 IstanbulS19 7.4 IstanbulS02, 7.4 IstanbulS12,
								7.4 IstanbulS20
Beijing, China	Shunyi–Qianmen– Liangxiang (TF)	6.5, 6.9, 7.3	70 × 70 × 30	200	1.5	Ideal scenarios	Antonietti <i>et al.</i> (2020)	Not included yet
Groningen, the Netherlands	NF	3.4	20 × 20 × 5	150	10	2018/01/08	Paolucci <i>et al.</i> (2021)	Not included yet
Kumamoto, Japan	Hinagu–Futagawa–Aso Caldera (SS)	7.0	53 × 46 × 22	500	1.5	2016/04/15 (16:25 UTC)	Sangaraju <i>et al.</i> (2021)	2016 7.0 Kumamoto
	Hinagu–Futagawa (SS) Aso Caldera (SS)	6.1 5.5				2016/04/14 (12:26 UTC) 2016/04/15 (18:03 UTC)		Not included yet Not included yet
The scenarios included in Bi chartwave velocity	B-SPEEDset are specified in the last co	lumn. f <sub>max</sub> , max	kimum frequency of the	e numerical r	nodel; <i>M</i> <sub>w</sub> , n	noment magnitude; NF, normal fa	ult; SoF, style of faulting; SS, strik	.e slip; TF, thrust fault, and $V_{{\sf S},{\sf min}}$ , minimum

TABLE 1



**Figure 1.** Workflow for postprocessing of SPEED results and creation of the corresponding flat file for BB-SPEEDset. Postprocessing codes are available to SPEED users on demand.

distributions is handled by means of specific MATLAB routines, implemented in the preprocessing tools of SPEED.

An overview of the postprocessing workflow is given in Figure 1, and it is organized in the following steps:

• 3PTOOL: The code extracts and organizes the raw SPEED seismograms in a common format (output file: Matlab .mat file), including receiver coordinates, displacement time histories (unfiltered), low frequency (directly from SPEED), and broadband (from Artificial Neural Network to BroadBand ground motions (ANN2BB); see Broadband Generation section) peak ground motion maps.

BB-SPEEDset trends with respect to those of the NESS dataset is presented, involving the statistical distribution of various intensity measures (IMs), their attenuation with distance, and the main features of near-source ground motion such as directionality, vertical components, and impulsive motions. Finally, an example of a query by earthquake scenario from the BB-SEEDset is introduced with reference to the 2009 L'Aquila earthquake, where one of the main advantages of a dataset based on PBS can be appreciated, in terms of generation of ground-motion maps of a selected earthquake scenario.

# WORKFLOW FOR POSTPROCESSING OF SPEED RESULTS

A fundamental step for the construction of a database of PBS results is the definition of an optimized workflow for postprocessing of results of SPEED simulations in a uniform and repeatable format. To this end, the SPEED kernel is supplemented by a set of MATLAB routine packages (available to SPEED users on demand) that allow to postprocess the raw waveforms computed by SPEED (typically, displacement time histories at receiver points), and generate outputs and metadata in a standard format. Note that, for the earthquake scenarios included in BB-SPEEDset, kinematic rupture models were introduced, consisting of heterogeneous slip functions across the fault. Although for real earthquakes (validations in Table 1), the kinematic source parameters were calibrated based on the available seismic source inversions studies, for scenario earthquakes the kinematic rupture generators proposed by Herrero and Bernard (1994) and Schmedes et al. (2012) are adopted. The generation of these fault-slip

- SITE RESPONSE PROXIES: The routine computes from the velocity model the most relevant site response proxies, namely:  $H_{bed}$ —depth of the alluvial-bedrock interface included in the simulation model,  $H_{800}$ —depth at which the shear-wave velocity  $V_s$  is equal or higher than 800 m/s,  $V_{s30}$ —time averaged shear-wave velocity from the surface to a depth of 30 m,  $V_{seq}$ —time averaged shear-wave velocity from the surface to  $H_{800}$  (if  $H_{800} \leq 30$  m; if  $H_{800} > 30$  m, then  $H_{800} = 30$  m), [see definition in the Italian Building Code (Norme Tecniche per le Costruzioni [NTC], 2018)],  $V_{sbed}$ -time averaged shear-wave velocity from the surface to  $H_{bed}$ ,  $V_{s800}$ —time averaged shear-wave velocity from the surface to  $H_{bed}$ ,  $V_{s800}$ —time averaged shear-wave velocity from the surface to  $H_{800}$ ; topography slope.
- EFFECTIVE FAULT: The code calculates the effective dimensions of the rupture fault area according to the procedure originally proposed by Mai and Beroza (2000), and extended by Thingbaijam and Mai (2016) (see Figure 2). This step is particularly relevant to define metadata with unbiased source dimensions, as the fault implemented in the numerical grid (typically, the fault associated with a maximum magnitude to be simulated) may be different from the coseismic rupture area associated with a given earthquake scenario. Effective source dimensions are based on the definition of autocorrelation width (Bracewell, 1986) of slip distributions, calculated in along-strike and downdip directions. These slip functions are computed summing up the slip in columns (or rows) on the rectangular rupture plane. An iterative, trimming process determines the largest dimensions that fit the autocorrelation width, according to the



**Figure 2.** Computation of the effective fault dimensions according to the procedure proposed by Mai and Beroza (2000) and extended by Thingbaijam and Mai (2016). The color version of this figure is available only in the electronic edition.

subfault size. The dependence of these effective measures on magnitude through scaling relationships has been verified (using e.g., Wells and Coppersmith, 1994 or Leonard, 2010).

- SELECT RECEIVERS: The code extracts subsets of receiver points according to prescribed sampling techniques. Because the simulated seismograms are generally obtained at tens of thousands of receivers, the computation of broadband time histories and of all corresponding IMs is limited to a subset of receivers to minimize the computational cost. In our processing for the BB-SPEEDset, the receiver selection was defined to achieve a higher density of receivers at lower distances from the source.
- BROADBAND GENERATION: At the selected receivers, the SPEED signals (typically reliable up to about 1.5–2 Hz; see Table 1) are enriched at high frequencies using a technique based on ANN trained on strong ground motion recordings (referred to as ANN2BB).
- SPEED IDCards: The routine produces an informative sheet (.pdf) summarizing the main features of the numerical model (e.g., mesh, wave velocity model), of the simulated fault rupture scenario (e.g., fault-slip distribution, rise time, rupture times), and of a selection of outputs (e.g., ground-shaking maps).
- FLAT FILE GENERATION: A flat file is created in a format consistent with the one adopted in up-to-date strong-motion databases (e.g., NESS) and populated with an exhaustive list of metadata (regarding the source, source-to-site distances, site response proxies, postprocessing, etc.) and ground-motion IMs.

#### **BROADBAND GENERATION**

As well known, the accuracy of the PBS is limited to the longperiod range  $T \ge T^*$ , with  $T^*$  being typically in the range of 0.5–1 s (see maximum frequency in Table 1), mainly due to the lack of knowledge about the Earth crust and earthquake rupture process at short wavelength and partially due to the computational cost of large and fine grids. For this reason, the core of the postprocessing workflow is the generation of broadband waveforms, in which the low-frequency simulated waveforms are enriched in the high-frequency range to produce time histories with a realistically broad frequency content. This is an essential step to treat the simulated waveforms in the same way as recordings, and, therefore, make them usable in earthquake engineering applications.

To this end, the ANN2BB

approach (Paolucci *et al.*, 2018), based on ANNs trained on strong ground motion recordings, is adopted. In this work, the ANN2BB technique has been enhanced with respect to the original version published in 2018 to make it suitable for massive postprocessing of larger datasets in a semiautomated fashion and to improve the quality of generated waveforms.

Referring to Paolucci *et al.* (2018) for further details, the ANN2BB procedure is based on four main steps, as sketched in Figure 3:

- An ANN, consisting of two-layer feed-forward neural network with 30 hidden neurons, is trained on a dataset of strong-motion records, such as SIMBAD v.6 (Smerzini *et al.*, 2014) or Next Generation Attenuation-West2 Project (Ancheta *et al.*, 2013), to predict short-period spectral ordinates (*T* < *T*\*) based on the long-period ones (*T* ≥ *T*\*). Different ANNs are trained separately on the geometric mean of the horizontal components and on the vertical components for prediction of three-component ground motions.
- 2. For each simulated waveform, a target ANN2BB response spectrum is computed, the long-spectral ordinates of which, for  $T \ge T^*$ , coincide with the simulated ones, whereas they are obtained from the ANN for  $T < T^*$  (separately for horizontal and vertical components). In our application, the corner period  $T^*$  is set depending on the frequency limitation of the numerical mesh. To preserve variability of the ANN results, the target at short periods is built on the median value calculated over 20 ANN realizations.
- 3. Once the ANN2BB target spectrum is defined for each waveform, the nonstationary stochastic approach by Sabetta and Pugliese (1996) is followed to generate the high-frequency portion of the signal. Updates of this approach (see e.g., Pousse *et al.*, 2006; Sabetta *et al.*, 2021) will be implemented in the next releases of the BB-SPEEDset. More specifically, 20



stochastic realizations are obtained, according to the simulated scenario  $(M_w)$  and receiver (distance, site conditions), and, out of these, a specific realization is selected based on the criteria illustrated in the sequel.

4. The selected stochastic signal (HF) and the PBS waveform (LF), previously filtered in the high- and low-frequency range, respectively, are combined in the time domain. Phase matching between HF and LF is achieved by alignment of the two time histories, according to the instant at 5% of normalized Arias intensity ( $I_A$ ).

The scaling and selection of the stochastic signal of step (3) is based on a two-step procedure of minimization of the residuals with respect to the target ANN2BB spectrum. First, for each stochastic realization, a scaling factor *SF* that minimizes the residuals with respect to the target is calculated, and, subsequently, among the different scaled stochastic signal the one with the minimum misfit is selected. The optimum realization is obtained using the following minimization procedure:

$$\min_{j=1:n \text{sim}} \left( \min_{SF} \sum_{i=1}^{N_{\text{periods}}} w_i \left( \frac{\ln \text{ANN}_{\text{median},i} - \ln(SF \times \text{STOCH}_i^j)}{\sigma_{\ln \text{ANN},i}} \right)^2 \right),$$
(1)

**Figure 3.** Flowchart of the ANN2BB approach revised after Paolucci *et al.* (2018) for the massive processing of physics-based numerical simulations (PBS) for broadband computation and compilation of BB-SPEEDset. The color version of this figure is available only in the electronic edition.

in which *SF* is the scaling factor (typically in the range 0.5–2),  $N_{\text{periods}}$  are the vibration periods of the target ANN spectrum,  $w_i$  the weight for the *i*-th period,  $n \sin = 20$  is the total number of stochastic signal realizations (20 realizations are found to guarantee a satisfying final spectrum). STOCH<sup>*j*</sup> is the *j*-th stochastic realization. Both median (ANN<sub>median</sub>) and standard deviation ( $\sigma_{\ln ANN}$ ) of the 20 ANNs are used to calculate the misfit. The weight vector *w* controls the fit to the target ANN spectrum. As the fit is actually performed only on the HF range, the weight vector takes the following values: w = 0 for  $T > T^*$ ; w = 1 for  $T < T^*$ , and w = 2 for  $T = T^*$ . These values were selected after appropriate sensitivity analyses, to ensure that the final spectrum is "as close as possible" to the PBS one in the LF range, including the merging period, at the end of the procedure.

The main differences with respect to the procedure introduced by Paolucci *et al.* (2018) are the following:

 step 2: The target ANN2BB in the high-frequency range is computed as the median of 20 different ANN realizations,

#### TABLE 2 Structure of the BB-SPEEDset flat file

Source Metadata	Scenario ID	Hypocenter lat/lon/depth	Strike, dip, rake	Fault mechanism
	Event ID	$M_{ m w}$	References	Rupture top
	Scenario ID card	M <sub>0</sub>	ANN_database	Fault vertices*
	Event_Time	Average slip	Transition period	Length*
	Event nation code	Number of segments	ANN2BB_procedure	Width*
Receiver Metadata	Receiver ID	-		
	Receiver east and nor	th coordinates		
	Receiver elevation			
Site Response Proxies	V <sub>530</sub>	<i>H</i> <sub>bed</sub> (depth of the alluvial-bedrock interface)	V <sub>Sbed</sub> (time averaged	$V_{\rm S}$ from the surface to $H_{\rm bed}$ )
	V <sub>Seq</sub>	$H_{800}$ (depth to $V_S \ge 800 \text{ m/s}$ )	V <sub>5800</sub> (time averaged	$V_{S}$ from the surface to $H_{800}$ )
Source-to-Site	Epicentral distance	<i>R_</i> line	Joyner and Boore dis	tance, R <sub>JB</sub>
Distances <sup>†</sup>			-	
	Hypocentral	Rx	<i>R</i> _rup	
	distance			
Intensity Measures <sup>‡</sup>	PGA	SA( <i>T</i> ) for <i>T</i> from 0.01 to 10 s	Pulse-like flag	Housner intensity
	PGV	Permanent displacement	Pulse period	Arias intensity
	PGD	Mean period	Ds595, Ds575	Cumulative absolute velocity

PGA, peak ground acceleration; PGD, peak ground displacement; and PGV, peak ground velocity.

\*Fault dimensions are given with respect to "numerical fault" and "effective fault.

<sup>†</sup>Distances from the fault are computed with respect to its "effective" dimensions.

\*Except pulse features, intensity measures are defined for the following directions: north–south, east–west, horizontal; up–down, vertical; HGM, horizontal geometrical mean; FN, fault normal; FP, fault parallel; D50, median value of the intensity measure (IM) distribution obtained from rotated waveforms; and D100, maximum value of the IM distribution obtained from rotated waveforms; (Rotation angles are given as well.)

rather than the output of a single ANN, to achieve a greater stability of results and minimize overfitting issues;

short periods, as it will be illustrated in the last section of this article for the case of the L'Aquila earthquake.

• step 3: The frequency scaling of the stochastic signal is replaced by a linear scaling in the time domain to approach the target ANN2BB spectrum at short periods. This improvement overcomes some issues related to the unrealistic frequency content that was sometimes introduced by the manipulation of Fourier spectra in the previous procedure. Furthermore, owing to its computational efficiency, it makes the broadband generation tool of SPEED suitable for massive computations over a very large number of receivers with no manual interventions, facilitating time-effective yet reliable outputs.

Finally, it is worth to remark two major advantages of the ANN2BB procedure with respect to other hybrid approaches to generate broadband signals based on PBS (see e.g., Mai and Beroza, 2003). First, since the resulting waveform is obtained based on the fit to a regular target response spectrum, it is possible to avoid the spurious discontinuities of the Fourier spectrum that are the typical result of the "glueing" of the physics-based (LF) signal with the stochastic one (HF), around a cross-over frequency. Second, as thoroughly discussed by Paolucci *et al.* (2018), in the ANN2BB procedure the LF and HF portions are not independent, as for most hybrid approaches, but their correlation is enforced by the records-trained ANN. This allows the BB-SPEEDset results to be used to produce realistic earthquake ground-motion maps with the proper spatial correlation also at

# OVERVIEW OF THE BB-SPEEDSET: METADATA AND IMs

For each scenario and each selected receiver, a list of source metadata, postprocessing metadata, receiver metadata, site response proxies, source-to-site distances, and IMs are computed and stored in a flat file (see Table 2 for details). The fields of the flat file are consistent with the ones in the engineering strong motion (Lanzano *et al.*, 2018) and the NESS flat files (Sgobba *et al.*, 2021). However, when considering a database of PBS results, it is relevant to store information regarding the type of postprocessing accomplished, such as ANN training database, ANN transition period, broadband procedure, as done in BB-SPEEDset.

Because directional effects may be significant in the nearsource region, each IM is defined on different horizontal directions, besides the vertical one, namely: the fault normal (FN) and fault parallel (FP) components, calculated rotating the horizontal waveforms orthogonal and parallel to the strike of the fault, respectively; the horizontal geometric mean (HGM), computed using the two horizontal east-west and north-south components; the maximum (RotD100) and the median (RotD50) values of IMs over all orientations (Boore, 2010), denoted in the following by D100 or D50, respectively. In addition to standard peak IMs, such as peak ground acceleration (PGA), peak ground velocity (PGV), peak ground



**Figure 4.**  $M_{\rm W}$  and  $R_{\rm JB}$  distribution of BB-SPEEDset (gray circles), in comparison with that from NEar-Source Strong-motion (NESS) dataset (black dots, after Sgobba *et al.*, 2021). The shaded region indicates the records excluded from the comparisons reported in this work.

displacement (PGD), and response spectral accelerations (SA), a variety of integral- and frequency-related IMs is included, such as the Housner intensity  $(I_{\rm H})$ , the cumulative absolute velocity (CAV), the Arias Intensity  $(I_A)$ , the  $I_A$  -based durations (i.e., time interval between 5% and 95% of the total  $I_A$ , Ds595, and between 5% and 75% of the total  $I_A$ , Ds575; see Trifunac and Brady, 1975; Bommer and Martinez-Pereira, 1999) and the mean period (Rathje et al., 1998). Furthermore, in the flat file compilation, special care was given to the characterization of pulse-like waveforms, which are of particular interest in earthquake engineering applications owing to their increased damage potential. Impulsive ground motions reflect two main physical effects. First, in forwarddirectivity conditions, the constructive superimposition of waves generated by a propagating rupture in front of a site may yield double-sided velocity pulses. Second, the contribution of waves generated by a finite dislocation on the fault plane can produce a permanent displacement (i.e., fling step), which results in a one-sided velocity pulse. As further discussed in the following, it is worth underlining that PBS can provide accurate predictions of displacement waveforms, including static offsets, which are hardly retrieved from recordings, because of the baseline drifts associated with errors in instrument response at low frequencies. The identification of pulse-like waveforms and of pulse period  $(T_p)$  has been done relying on the algorithm proposed by Shahi and Baker (2014). A thorough discussion on impulsive ground motions will be



**Figure 5.** Distribution of BB-SPEEDset with respect to  $V_{S30}$ . The color version of this figure is available only in the electronic edition.

provided later in the article (see the Analysis of impulsive ground motions section).

The distribution of BB-SPEEDset data with respect to magnitude and distance is given in Figure 4. The dataset includes a total of 12,058 three-component waveforms from earthquake scenarios with  $M_w$  from 5.5 to 7.4 and Joyner–Boore distances ( $R_{\rm JB}$ ) up to 80 km. Strike-slip, normal, and thrust events are included in the dataset. Most records refer to normal (50%) and strike-slip (41%) focal mechanisms, whereas only 9% is from thrust earthquakes (i.e., only the 2012 Po Plain event; see Table 1). The dominance of normal and strike-slip faults is because, on one hand, normal events are typical of the seismicity in central Italy (mostly represented within the BB-SPEEDset), and, on the other hand, a significant set of strike-slip events from the North Anatolian fault (Istanbul case study, see Infantino *et al.*, 2021) and from the 2016 Kumamoto sequence (Sangaraju *et al.*, 2021) is also included.

In the same figure, the distribution of the NESS dataset is illustrated for comparison. Although the BB-SPEEDset does not cover yet earthquakes with  $M_{\rm w} \geq 7.5$  (and, for this reason, the comparisons shown hereafter will neglect this range), it can be seen that PBS allow to approach an ideally dense sampling at short distances that cannot be obtained by the NESS recording stations.

The distribution of BB-SPEEDset with respect to site conditions, parametrized in terms of the  $V_{S30}$ , is shown in Figure 5. About 40% of data correspond to soil conditions with  $V_{S30} < 800$  m/s and about 60% to rock conditions, with dominance for hard-rock sites with  $V_{S30}$  larger than 1000 m/s. Within the soil classes, the majority of waveforms is on stiff soil with  $V_{S30} > 400$  m/s, but an appreciable number of data are on very soft sites with  $V_{S30}$  as low as 150 m/s (Marsica case study). Figure 5 highlights another potential advantage of the simulated datasets: seismostratigraphic conditions are fully known in the PBS, so that each receiver may easily be associated to different proxies related to site response. Furthermore, site conditions that

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**Figure 6.** Cumulative distribution functions of peak ground acceleration (PGA), peak ground velocity (PGV), spectral acceleration (SA)(1.0 s), SA(3.0 s), cumulative absolute velocity (CAV), Ds595, Arias intensity ( $I_A$ ), and Housner intensity ( $I_H$ ), as obtained from BB-SPEEDset (dark gray: empirical; black: best-fitting lognormal distribution with corresponding statistical moments) and from NESS (light gray and dashed lines). The 95th percentiles of the statistical distributions are also superimposed on the graph. For all intensity measures (IMs), the D50 component is considered.

increase in a short time by extending to the case studies listed in Table 1 and not processed yet. Furthermore, new scenarios may specifically be developed to fill in the magnitude, distance, and site conditions gaps of the present version of BB-SPEEDset.

# STATISTICAL DISTRIBUTIONS OF GROUND-MOTION IMs

As remarked in the Introduction, a crucial step for the potential use of the BB-SPEEDset flat file and signals for engineering applications is to compare the statistical distributions of different IMs, their attenuation with distance, the features of directional and impulsive near-source accelerograms, with those obtained from recorded ground motions, to verify the presence of potential biases and to identify their sources. For this purpose, because the SPEED results essentially refer to near-source conditions, we considered the NESS dataset introduced previously as a reference.

Figure 6 shows the cumulative distribution function of different IMs, namely, PGA, PGV, SA(1.0 s), SA(3.0 s), CAV, Ds595,  $I_A$ , and  $I_H$ , as computed from the entire BB-SPEEDset flat file. For all IMs, the D50 component is considered. To verify the consistency of our results against recordings, the statistical distributions derived from BB-SPEEDset are compared in Figure 6 with those

are typically poorly represented in the recorded datasets, such as rock and very soft sites, may be better sampled. This may provide further constraints, complementary to those from earthquake recordings, for the calibration of site amplification factors, especially in complex geological conditions.

As a final remark, BB-SPEEDset should be considered as a dynamically growing dataset, the scenarios of which may

obtained on the NESS dataset, within similar ranges of  $M_w$ and distances. As a matter of fact, for this comparison, only records with  $M_w < 7.5$  are considered (corresponding to about 55% of NESS, see Fig. 4), consistently with BB-SPEEDset; note that, at low magnitude, larger distances are covered by BB-SPEEDset. To emphasize the comparison between the two independent sets of data, the lognormal distributions, fitting



**Figure 7.** SA(0.1 s) and SA(5 s) distribution from BB-SPEEDset (dark gray) in comparison with that from NESS (light gray) in the same range of magnitude. The D50 component is considered for both spectral accelerations.

the empirical ones, are also superimposed, and the corresponding statistical moments are provided. Furthermore, for each IM distribution, the values of ground motion with a probability of exceedance less than 5% (i.e., 95th percentile of the related distributions) are highlighted on the graph. Overall, it is remarkable that a full consistency is found between BB-SPEEDset and NESS, as it should be, as the distribution of  $M_w$  and  $R_{JB}$  of the two datasets are also consistent (apart from slight differences, as



**Figure 8.** D50 components of (a–c) PGA and (d–f) SA (1 s) (bottom) as a function of  $R_{JB}$  distance, for BB-SPEEDset (dark gray) and NESS (light gray), considering different  $M_{W}$  ranges (centered around 6.0, 6.5, and 7.0).

commented earlier). Although the BB-SPEEDset tends to underestimate the recorded values of PGA by a factor of about 20%, probably owing to the difficulty of the ANN2BB approach to describe the short periods as accurately as the long periods, the statistical distributions of several IMs, from peak measures to integral ones, are noticeably similar both in terms of median and standard deviation ( $\sigma_{\log 10}$ ). The agreement of PGV, SA (1.0 s), CAV, and  $I_{\rm H}$  is excellent. For SA(3.0 s), BB-SPEEDset tends to provide more severe ground motions, most likely because of the intrinsic higher coherency of waveforms simulated by means of numerical models that inevitably cannot account for the actual small-scale heterogeneities and complexities in the source, path, and site. Referring to  $I_A$ , some discrepancies are found especially at intensity values lower than the median values (median from BB-SPEEDset is lower than NESS of a factor of about 25%), but the agreement improves significantly above the median. As a matter of fact, the 95th percentiles from BB-SPEEDset and NESS differ of less than 10%. Although the standard deviations from BB-SPEEDset and NESS are comparable in most cases, differences are found for the duration Ds595, being  $\sigma_{\log 10}$  from BB-SPEEDset lower (0.18) with respect to the one from NESS (0.25), suggesting that the level of waveform complexity achieved through SPEED simulations is still lower than reality.

As a further consistency check, the distribution of SA(0.1 s) and SA(5.0 s) obtained from BB-SPEEDset is compared with that from NESS ( $M_w < 7.5$ ) in Figure 7. It turns out that the correlation between long and short periods of BB-SPEEDset (the latter ones being a direct output of the ANN2BB procedure described previously) is consistent with the one from

NESS, with few exceptions of NESS values having a combination of high short-period and low long-period spectral ordinates, not present in the simulated waveforms.

## GROUND-MOTION ATTENUATION WITH DISTANCE

Attenuation of ground-motion IMs with distance for the BB-SPEEDset and NESS are compared in this section. Figure 8 shows PGA and SA(1 s) as a function of  $R_{\rm JB}$ , using D50 component, for different  $M_{\rm w}$  ranges. No discrimination of soil conditions has been made. The agreement is reasonably good, with a similar trend of IMs with distance. As already mentioned, there is a tendency of PGAs of the BB-SPEEDset to lie on

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**Figure 9.** Duration Ds595 (for D50 component) versus  $R_{JB}$  distance, for both BB-SPEEDset (dark gray) and NESS (light gray), for different  $M_{w}$  ranges (centered around 6.0, 6.5, and 7.0).



**Figure 10.** Ratios of (a–c) fault normal and fault parallel (FN/FP) and of (d–f) V/D50 for PGA (left), PGV (center), and SA (3 s) (right), as a function of  $R_{JB}$ . The median  $\pm \sigma$  ratios from BB-SPEEDset (dark gray) and NESS (light gray) are compared.

1994 Northridge earthquake, whereas the lowest values (in the same  $M_w$  and  $R_{JB}$  ranges) are for the 2014 Aegean Sea earthquake of  $M_w$  6.4. The very low SA(1 s) values for  $M_w$  in the 6.8–7.2 range and  $R_{JB} > 20$  km are related to the Japanese earthquake of 2008.

Figure 9 shows attenuation with  $R_{\rm JB}$  for the Ds595 duration of ground motion. Again, an overall good agreement among the two datasets is noticeable, with NESS records showing a greater scatter probably because of their higher complexity than simulated waveforms, as previously noted. Except for this remark, simulated durations provide a consistent trend with magnitude and distance.

## DIRECTIONALITY AND VERTICAL-TO-HORIZONTAL MOTIONS

It is widely recognized that earthquake ground motion may exhibit specific features in proximity of the source (e.g., Stewart et al., 2001), including polarization related to the fault mechanism and large, shortperiod, vertical components exceeding, even significantly, the corresponding horizontal ones (Bommer et al., 2011; Gülerce and Abrahamson, 2011). These aspects will be addressed in this section, with special care again to the comparison with recordings.

igure 10 shows the median

the lower side of recorded values, especially in the lower  $M_w$  range. Spectral ordinates at intermediate periods, SA(1 s), show a better agreement with NESS, at any distance. A large but comparable scatter can be observed for the two datasets, slightly larger for SA(1 s). The larger scatter for  $M_w$  around 7.0 and  $R_{\rm JB} > 10$  km is related to the Istanbul simulations that do not have receivers at shorter distances from the source. The high spectral accelerations from NESS at 1 s for  $M_w$  in the 6.3–6.7 range and  $R_{\rm JB} > 10$  km are mostly related to the

 $(\pm \sigma, \text{ shaded regions})$  FN/FP (Fig. 10a–c) and V/D50 (Fig. 10d–f) for PGA, PGV, and SA(3 s), as a function of  $R_{\text{JB}}$ , as obtained from BB-SPEEDset and NESS ( $M_{\text{w}} < 7.5$ ). Although long-period components of ground motion are directly related to the PBS results, short-period components, such as PGA, reflect the output of the ANN2BB procedure. Referring to the FN/FP distribution, an excellent agreement is found both in terms of median value and variability at different distances from the source. As expected, both datasets show



**Figure 11.** Pulse period  $T_p$  versus earthquake magnitude for observed pulselike ground motions of BB-SPEEDset (20% of set data, in light gray) and NESS dataset (almost 30% of whole set data, in black).

that in near-source conditions the FN motion is stronger than the FP at intermediate and long periods and at very short distance,  $R_{\rm JB} < 5$  km. At increasing distances, the FN polarization tends to vanish, in agreement with previous studies (Somerville *et al.*, 1997; Pacor *et al.*, 2018). On the other hand, at short periods, no polarization effects are found, with median FN/ FP ratios equal to approximately one at all distances, both in BB-SPEEDset and in NESS. Especially at long periods, the variability across periods is rather large, meaning that directionality features are region- and scenario-specific. Here, the variability of simulated waveforms tends to be slightly larger, but differences remain limited. Further insights on the physical reasons of such variability could be obtained by analyzing the dependence on the focal mechanism and the spatial distribution of such ratios for specific rupture scenarios.

As regards to the V/D50 ratios, there is a good agreement in terms of PGA (median V/D50 ~ 0.6 for both datasets, but with larger variability for NESS), whereas for PGV and SA(3 s) the BB-SPEEDset values at very short distance (median V/D50 ~ 0.6) are higher than NESS (median V/D50 ~ 0.4). More detailed studies are planned to investigate the dependence of such different ratios on the focal mechanism and site conditions.

### **ANALYSIS OF IMPULSIVE GROUND MOTIONS**

This section focuses on the identification and comparison of pulse signals, together with the corresponding period  $T_p$ , tagged and stored in both BB-SPEEDset and NESS flat files.

Figure 11 shows  $T_p$  values calculated following the algorithm proposed by Shahi and Baker (2014) as a function of

magnitude, for both BB-SPEEDset and NESS.  $T_p$  ranges from 1 to 12 s, with amplitudes tending to increase with magnitude (Mavroeidis and Papageorgiou, 2003; Somerville, 2003), but with the largest values probably related to coupling with deep basin conditions, such as for the earthquakes of  $M_w$  6.0 Po Plain, and  $M_w$  6.7 Marsica, that are poorly represented by the NESS dataset. With these limitations in mind in terms of comparison of datasets, the agreement of the  $T_p$  trends from the two datasets is remarkable.

Figure 12 shows the trend of  $T_p$  as a function of the ratio between PGD and PGV (*PGD*/*PGV*), for NESS (Fig. 12a) and BB-SPEEDset (Fig. 12b). A reasonably good agreement between the two sets is found within comparable magnitude ranges (i.e.,  $M_w$  from 6.0 to 7.4). To improve the accuracy of the PGD estimations from the NESS dataset,  $T_p$  and peak values from the e-BASCO baseline corrected waveforms have been considered (see Sgobba *et al.*, 2021 for details).

Closed-form analytical relationships between  $T_p$  and PGD/PGV are also shown in Figure 12, with thick black and gray lines, calculated based on simple functions that may approximate impulsive ground motion. These relationships are based on the analytical expressions of the Fourier spectra of the "Ricker wavelet" and of the "double-impulse" functional forms, given by (for unit peak amplitude):

Ricker wavelet : 
$$v(t) = (1 - 2\alpha^2 t^2)e^{-\alpha^2 t^2}$$
 with  $\alpha = \pi f_p$ ,  
(2a)

Double-impulse : 
$$v(t) = 2\pi f_p \sqrt{e}t e^{-\alpha^2 t^2}$$
 with  $\alpha = \sqrt{2}\pi f_p$ ,  
(2b)

for which the peak of the Fourier spectrum corresponding to the frequency  $f_p = 1/T_p$  can be identified and related to PGV and PGD, as shown in the top right of Figure 12. It is interesting that most data from both sets fall between the two analytical relationships, suggesting that the variability of impulsive ground motions may roughly be represented by these two functional families.

# QUERY BY EARTHQUAKE OF THE BB-SPEEDset: THE L'AQUILA EXAMPLE

In this section, we highlight the information that can be obtained from BB-SPEEDset by selecting a specific earthquake (query by earthquake) and extracting the corresponding waveforms and IMs. As an explanatory case study, the 6 April 2009  $M_{\rm w}$  6.2 L'Aquila earthquake is considered.

First, as a set of illustrative results of the ANN2BB procedure, Figure 13 shows the broadband three-component (east-west, north-south, and up-down) acceleration, velocity, and displacement time histories at selected receivers, superimposed on the corresponding PGA, PGV, and PGD maps. Referring to

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**Figure 12.** Pulse period  $T_p$  as a function of *PGD/PGV* for (a) NESS and for (b) BB-SPEEDset, for different magnitude ranges, together with analytical relationships for Ricker wavelet and double-impulse functions.

Smerzini and Villani (2012) and Evangelista *et al.* (2017) for a thorough comparison of recorded and simulated waveforms, it is noted herein that simulated time histories have realistic features in terms of duration, amplitudes, and frequency content, with displacement waveforms showing permanent displacements related to the coseismic slip on the fault. Furthermore, the spatial distribution of peak ground motion values at high frequency (PGA) turns out to be well correlated with the geological features of the basin (see areas of maximum amplitudes within the basin), supporting the effectiveness of the ANN2BB approach in establishing a correlation between long- and short-period ordinates, and, thus, reproducing at short periods physics-based features that are simulated only at long periods. It is also interesting

to note that, moving from maps related to short periods (PGA) to those related to long periods (PGD), the spatial correlation of ground motion appears to be characterized by increasing correlation length, as expected, at least from a qualitative point of view. Quantitative evaluations of spatial correlation of accelerations spectral from broadband numerical simulations can be found in Infantino, Smerzini, and Lin (2021) and in Schiappapietra and Smerzini (2021).

To appreciate the richness of information included in

BB-SPEEDset flat file, Figure 14 shows, for the same case study of L'Aquila, the maps of different IMs, namely, PGV, CAV,  $I_A$ ,  $T_p$ ,  $I_H$ , and Ds595. The FN component is shown for all IMs, except for  $T_p$ , which is computed for the orientation in which the pulse is the strongest according to the algorithm by Shahi and Baker (2014).

One of the main outcomes of BB-SPEEDset is related to the possibility of drawing maps and making quantitative evaluations of spatial correlation of a broad spectrum of IMs, for any orientation of ground motion, with a level of detail that could not be possible using recordings owing to their limited number. From the maps of Figure 14, the following comments can be made:





whereas the dashed contour denotes the location of the L'Aquila historical center. The inset in (a) shows the location of the map within Italy. The color version of this figure is available only in the electronic edition.



- areas of maximum intensity of PGV, CAV, *I*<sub>H</sub>, and *I*<sub>A</sub> are concentrated on the surface projection of the fault and inside the alluvial basin, because of the coupling of site response with rupture propagation effects;
- lobes of maximum CAV correlate well with the shape of the basin, also at its southeastern edge;
- there is an overall similarity between the spatial patterns of PGV and  $I_{\rm H}$ , which is not obvious from their definitions: the former is an instantaneous measure of ground shaking, whereas the latter, defined as the area of the pseudovelocity spectrum (5% damping in this case) between 0.1 and 2.5 s, is proportional to the maximum kinetic energy stored in an elastic structure;
- the areas of pulse-like ground-motion occurrence are found typically close the top edge of the fault, mostly related to up-dip directivity effects, highlighting an interesting correlation between strong pulse-like motions with the areas of peak values for the FN component (see left panels of Fig. 14.);
- correspondingly to areas with impulsive motions, the lowest values of ground-motion duration are found.

Although these comments may be considered to be specific for the L'Aquila earthquake case study, they convey the idea

**Figure 14.** 2009  $M_w$  6.2 L'Aquila: maps of different ground motion IMs stored in BB-SPEEDset, (a) PGV, (b) CAV, (c)  $I_A$ , (d) pulse period  $T_p$ , (e)  $I_H$ , and (f) DS595. For all IMs, except  $T_p$ , the FN component is shown. The color version of this figure is available only in the electronic edition.

of the comprehensive picture of earthquake ground motion that can be obtained from PBS and of the potential outcomes that the analysis of the simulated waveforms may provide in terms of an improved characterization of near-source ground shaking. As a further example, Figure 15 shows the maps of permanent displacement  $D_{\text{perm}}$  computed from BB-SPEEDset for the L'Aquila case study, along the FP, FN, and UD components. Although the fling step is naturally reproduced by physics-based simulation of seismic-wave propagation, the estimation of permanent displacement from earthquake recordings requires complex signal processing procedures (D'Amico et al., 2019), which are typically subject to high uncertainties. This points out one of the main advantages of numerical simulations, that is, to provide an accurate and detailed picture of the long-period components of earthquake ground motion, with relevant implications for the calibration



of GMMs for PGD and displacement spectral ordinates (Cauzzi et al., 2015).

Coherently with the normal focal mechanism of the Paganica fault, responsible of the L'Aquila earthquake, the ground underwent a static maximum horizontal offset of around 10 cm (both on the footwall and hanging wall along the FN and FP directions, respectively) and a maximum subsidence of about 15 cm on the hanging wall. The latter is consistent with the maximum coseismic vertical displacement obtained from both Differential Interferometric Synthetic Aperture Radar and Global Positioning System observations (D'Agostino *et al.*, 2012).

### CONCLUSIONS

In this article, we have presented the BB-SPEEDset—a new dataset of near-source broadband earthquake ground motions from 3D PBS obtained by the spectral element computer code SPEED. This is expected to support research on the characterization of earthquake ground motions in the proximity of the seismic source and in complex geological conditions that cannot be extensively documented based on available records.

To produce the dataset, an effective workflow has been devised to postprocess raw PBS results in a homogeneous and repeatable format. The core of the workflow is the generation of broadband ground motions starting from PBS results, reliable only in the low-frequency range, according to the ANN2BB procedure first proposed in Paolucci *et al.* (2018) and further improved in this work to make it suitable for massive processing of simulated waveforms.

At the present stage, BB-SPEEDset consists of a total number of 12,058 three-component waveforms from worldwide earthquake scenarios, mostly validated against records, with  $M_{\rm w}$  from 5.5 to 7.4 and  $R_{\rm JB}$  up to 80 km. Besides source, receiver, and postprocessing metadata, the BB-SPEEDset flat file can provide a large portfolio of ground-motion IMs, from the standard peak and spectral measures to integral ones (e.g.,  $I_{\rm H}$ , CAV, duration,  $I_{\rm A}$ ), up to parameters related to impulsive

**Figure 15.** 2009  $M_{\rm w}$  6.2 L'Aquila: maps of permanent displacement ( $D_{\rm perm}$ ) on the three components of motion, (a) FP, (b) FN, and (c) up–down (UD). The color version of this figure is available only in the electronic edition.

ground motions (e.g., pulse period  $T_p$ ) and long-period components of ground motions such as  $D_{perm}$ .

An extensive set of checks has been performed and documented in this article to verify that the BB-SPEEDset provides peak values, integral IMs, ratios of long-to-short period spectral ordinates, features of impulsive ground motions and directionality effects, consistent on a statistical basis with NESS-a dataset of worldwide near-source records (Sgobba et al., 2021). The positive outcome of such consistency check was not obvious, because the BB-SPEEDset is the last step of a series of complex studies, starting from the construction of large-scale 3D numerical models, the simulation of realistic fault-rupture scenarios, the source-to-site propagation of seismic waves in complex geological media, and, finally, a smart postprocessing of low-frequency signals to get broadband waveforms using ANN. Although reproducing exactly the recorded ground motions by PBS is not an achievable objective, this article shows that it is possible to construct realistic earthquake ground-motion scenarios, and that the resulting waveforms are consistent in terms of peak values, duration and frequency content, with records obtained in near-source conditions.

Given this major outcome, we envisage that the BB-SPEEDset, either in the present version or in the following ones enriched by further simulated scenarios, will serve as the basis for several new achievements for an improved characterization and engineering usage of near-source earthquake ground motions, such as:

 to fill in the gaps, in terms of source-to-site conditions, focal mechanisms, variability of fault-slip distributions and directivity effects, complex geological conditions that are present in the worldwide near-source records datasets and that are not expected to be easily covered in short time by additional records;

- in the conditions above, provide region- and sitespecific input motions for nonlinear structural analyses of engineered structures that are presently often carried out using unrealistic scaling factors on recorded ground motions;
- to provide accurate predictions of long-period components of ground motions, including PGDs and static offsets, that are hardly retrieved from records because of the uncertainties associated with the postprocessing procedures;
- to construct region-specific scenarios of earthquake ground shaking, suitable to improve empirical models of spatial correlation (Infantino, Smerzini, and Lin, 2021) and spatial coherency of ground motion (Smerzini, 2018), taking advantage of the dense spacing of receivers that can be achieved in the numerical modeling;
- to support the development of nonergodic models for ground-shaking scenarios as a key tool for enhanced seismic hazard and seismic risk evaluations in large urban areas (Stupazzini *et al.*, 2021).

#### **DATA AND RESOURCES**

The NEar-Source Strong-motion version 2.0 (NESS2.0) dataset has been downloaded at http://ness.mi.ingv.it/ (last accessed March 2021). The BB-SPEEDset (v.1.0) is available at http:// speed.mox.polimi.it/BB-SPEEDset (last accessed May 2021), in which both the flat file and corresponding broadband waveforms can be downloaded. The open-source Spectral Element code SPEED is available at http://speed.mox.polimi.it (last accessed July 2021).

# **DECLARATION OF COMPETING INTERESTS**

The authors acknowledge that there are no conflicts of interest recorded.

## ACKNOWLEDGMENTS

This work has been partially supported by swissnuclear within the research activity "Development of advanced numerical approaches for earthquake ground motion prediction," in the framework of the SIGMA2 project, and by the Department of Civil Protection within the ReLUIS project WP18 "Normative contributions related to seismic action." The authors wish to thank in particular, Marco Stupazzini, Ilario Mazzieri, Maria Infantino, and Karim Tarbali, together with the entire SPEED team, for their support in the development of SPEED kernel, of the related pre- and postprocessing tools and for their contributions to the simulation of some of the earthquake scenarios that this work has been based on. The fruitful discussions with Philippe Renault in the framework of the swissnuclear-SIGMA2 project are also gratefully acknowledged. Constructive comments by Thomas Pratt, Francesca Pacor, and by another anonymous reviewer helped improving the article and are gratefully acknowledged.

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Manuscript received 22 March 2021 Published online 20 July 2021